

Animation of Open Water Phenomena with coupled Shallow Water and Free Surface Simulations

Nils Thürey¹, Ulrich Rüdè¹ and Marc Stamminger²

¹ Institute for System Simulation, University of Erlangen-Nuremberg, Germany

² Institute for Computer Graphics, University of Erlangen-Nuremberg, Germany

Abstract

The goal of this paper is to perform simulations that capture fluid effects from small drops up to the propagation of large waves. To achieve this, we present a hybrid simulation method, that couples a two-dimensional shallow water simulation with a full three-dimensional free surface fluid simulation. We explain the approximations imposed by the shallow water model, and how to parametrize it according to the parameters of a 3D simulation. Each simulation is used to initialize double layered boundary conditions for the other one. The area covered by the 2D region can be an order of magnitude larger than the 3D region without significantly affecting the overall computation time. The 3D region can furthermore be easily moved within the 2D region during the course of the simulation. To achieve realistic results we combine our simulation method with a physically based model to generate and animate drops. For their generation we make use of the fluid turbulence model, and animate them with a simplified drag calculation. This allows simulations with relatively low resolutions.

Categories and Subject Descriptors (according to ACM CCS): I.3.7 [Computing Methodologies]: Computer GraphicsAnimation; I.6.3 [Computing Methodologies]: Simulation and ModelingApplications;

1. Introduction and Related Work

In previous years, the physically based animation of liquids has seen significant progress. Especially the work of Ron Fedkiw [FAMO99], his group [LGF04] [IGLF06], and Jos Stam [Sta99] have enabled the realistic and efficient simulation of liquids with a free surface. A typical animation to show the quality of a simulation is pouring liquid into a glass. Nowadays the simulations recreating these effects are hard to distinguish from reality. In this paper we will, however, focus on scenes that have a larger scale, e.g. a ship traveling through the ocean. The challenge is to capture both the large scale movement of the water around the ship, as well as the splashing of the waves around it, including drops and spray. As all these phenomena contribute to the visual appearance, they have to be captured to achieve a realistic representation of such a scene.

For all classes of algorithms that are used to simulate free surface flows, the problem is that the amount of computational work and the required resources grow significantly when the resolution of the simulation is increased. The full

simulation of the above mentioned scene with a volume-of-fluid Navier-Stokes solver would hardly be possible even on large supercomputers. Adaptive techniques can be used to alleviate this problem to some extent [LGF04], but usually increase the complexity of a solver and have limits in their ability to speed up the computational time. In the following, we will describe a different approach that computes the full fluid flow only in a bounded region of interest, and uses a fast two-dimensional fluid simulation to compute the fluid surface around it. We also choose not to simulate the whole depth of the fluid, from the free surface to the bottom (e.g. the ocean floor), but only an upper layer of fluid. The small scale details such as drops are simulated as particles with a simplified, yet physically based, algorithm. The contributions of this paper are:

- a new hybrid method that couples a 2D shallow water simulation with a full 3D free surface simulation, and
- a physically based model for the creation and simulation of drops.

Although we use the lattice Boltzmann method to solve the two- and three-dimensional fluid flows, variants of both our hybrid method and the drop model could also be applied to other types of fluid solvers.

In its different forms the *Navier-Stokes* (NS) equations have long been used for physically based animation. [KM90] were the first to use *shallow water simulations* in computer graphics. As of today, this simplified model, which assumes depth averaged fluid properties, is still a research topic e.g. for computations performed on the GPU [HHL*05]. The use of a three-dimensional NS discretization with free surface boundary conditions was demonstrated by [FM96]. By today, various methods exist to directly solve or approximate the NS equations. Apart from the level set based solvers mentioned above, *volume-of-fluid* (VOF) methods are commonly used [MMS04], or, from the background of astrophysics, the *smoothed particle hydrodynamics* (SPH) solvers [KAG*05]. To model large scale water surfaces deep-water wave models are often used to compute the water surface motion, e.g. in [Tes04] or [HNC02]. These, however, do not solve the NS equations directly, but model the wave propagation by different forms of trochoid function spectra. In [RNGF03], Rasmussen et al. combined a 2D flow field with periodic 3D simulations to create detailed smoke simulations. Large scale open water scenes with 2D and 3D simulations were also used in TV and feature film productions, however, without giving detailed information about the models used. Recently, an approach to optimize a fluid simulations with tall and thin cells was proposed by [IGLF06], thus also reducing the computational complexity for large fluid volumes.

In the following we will use the *lattice Boltzmann method* (LBM), which originates from discrete compressible gas simulations [FdH*87]. The LBM approximates the NS equations without the need for an iterative solver by relaxing the incompressibility constraint [HL97]. It has been used in various areas, e.g. for single phase fluid computations [WZF*03], or in combination with a VOF model to simulate free surface flows [TR04]. They also explain how typical limitations of VOF methods can be overcome within the LBM algorithm. As the shallow water equations represent an advection-diffusion problem similar to the full NS equations, they can likewise be solved with the LBM. A derivation of the appropriate changes to the basic algorithm can be found in [Del01].

A different approach for detailed and accurate fluid simulation solvers, that we will make use of in Section (4), can be found in the area of chemical engineering. For cases such as bubble column reactors, Eulerian-Lagrangian simulations of these dispersed multi-phase flows, e.g. large numbers of bubbles in a relatively coarse fluid flow simulation [DKvS99], are used to understand and optimize the physical processes [BGD05]. These methods simulate bubbles with a spherical shape, and model the forces caused by the turbulent fluid

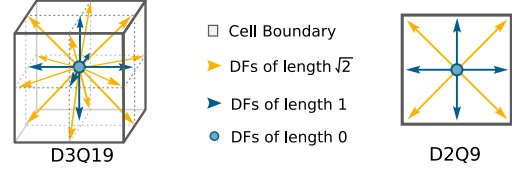


Figure 1: The lattice Boltzmann grid models used for the free surface ($D3Q19$) and the shallow water simulations ($D2Q9$).

around them. Apart from level set methods, where particles are used to accurately track the free surface, [TFK*03] also use particles to add small scale details. But in contrast to our approach they generate drops based of the surface curvature, and apply linear damping to model air resistance.

2. Animation of the Water Surface

Within a certain region of interest, e.g. around a moving ship, we perform a full three-dimensional simulation of the free surface fluid. We will briefly describe the algorithms used to perform the three- and two-dimensional fluid simulations, and then describe the coupling of both in more detail.

Free Surface Simulation: For solving the fluid phase in the three dimensional region we use the $D3Q19$ lattice Boltzmann model shown in Figure 1. This model requires an equidistant grid (with a cell size of Δx) and 19 floating point values, one for each velocity vector, that are stored in each cell. These *distribution functions* (DFs) represent a small amount of fluid moving with the corresponding velocity. The basic algorithm proceeds in two steps - first the advection of the fluid molecules is handled by copying the DFs to the neighboring cell along their velocity direction. Once this is done for all cells in the grid, each again has a full set of DFs, which is used to calculate the macroscopic properties of the fluid - the density ρ and the velocity \mathbf{u} . In the following, f_i will denote one of the nineteen DFs, with \mathbf{e}_i being the corresponding velocity vector. Its inverse direction is $\mathbf{e}_{\bar{i}}$, thus $\mathbf{e}_{\bar{i}} = -\mathbf{e}_i$.

$$\rho = \sum f_i \quad \mathbf{u} = \sum \mathbf{e}_i f_i . \quad (1)$$

Density and velocity are needed to calculate the equilibrium DFs with

$$f_i^{eq}(\rho, \mathbf{u}) = w_i \left[\rho + 3\mathbf{e}_i \cdot \mathbf{u} - \frac{3}{2}\mathbf{u}^2 + \frac{9}{2}(\mathbf{e}_i \cdot \mathbf{u})^2 \right] . \quad (2)$$

Here w is a weight that depends on the length of the velocity vector: $w_i = 1/3$ for $i = 1$, $w_i = 1/18$ for $i = 2, \dots, 7$, and $w_i = 1/36$ for $i = 8, \dots, 19$. The equilibrium DFs from Eq. (2) are necessary to compute the collisions that occur between the molecules in a real fluid. With the LBM these are calculated by relaxing the DFs towards the equilibrium DFs depending on the fluid viscosity with:

$$f_i(\mathbf{x}, t + \Delta t) = (1 - \omega)f_i'(\mathbf{x}, t + \Delta t) + \omega f_i^{eq} . \quad (3)$$

Here the relaxation parameter ω is calculated from the time step of the simulation Δt , the physical viscosity ν' and the lattice viscosity ν with

$$\omega = \frac{1}{3\nu} + \frac{1}{2}, \nu = \nu' \frac{\Delta t}{\Delta x^2}. \quad (4)$$

The free surface model we apply only simulates the fluid phase. Hence, DFs that would be streamed from the gas phase have to be reconstructed in cells at the interface with

$$f'_i(\mathbf{x}, t + \Delta t) = f_i^{eq}(\rho_A, \mathbf{u}) + f_i^{eq}(\rho_A, \mathbf{u}) - f_i(\mathbf{x}, t), \quad (5)$$

where ρ_A is the reference density of the air, hence, in our case $\rho_A = 1$. These free surface boundary conditions do not include surface tension, but ensure an undisturbed movement of the fluid in the gas phase. The mass flux between cells is computed directly from the DFs that are streamed from and to interface cells. This mass value m is additionally stored for each cell, and can be used to determine the fluid fraction $\varepsilon = m/\rho$ of a cell. Thus, this value determines how much of a cell is filled with fluid. Further details for a free surface implementation with LBM can be found in e.g. [KPR*05]. At the boundary of the 3D domain we use boundary conditions with a height dependent pressure and a velocity given by the shallow water simulation. In Section (3) this will be described in more detail.

Turbulence Model: A problem of the basic LBM is, that it can exhibit instabilities for lower viscosities and thus values of ω close to 2. To alleviate this, turbulence models are often applied. We make use of the commonly used *Smagorinsky* turbulence model [Sma63], that locally increases the fluid viscosity in regions where unresolved flow features are detected, to model the energy dissipation of small scale vortices [HSCD96]. The local increase of viscosity is determined by the Reynolds stress tensor Π , that is computed locally for each cell as

$$\Pi_{\alpha,\beta} = \sum_{i=1}^{19} \mathbf{e}_{i\alpha} \mathbf{e}_{i\beta} (f_i - f_i^{eq}), \quad (6)$$

This value is used to compute the correction factor

$$S = \frac{1}{6C^2} \left(\sqrt{\nu^2 + 18C^2 \sqrt{\Pi_{\alpha,\beta} \Pi_{\alpha,\beta}}} - \nu \right). \quad (7)$$

with a user defined constant C , that we have set to 0.04. The modified viscosity of the cell is then calculated with

$$\omega_s = \frac{1}{3(\nu + C^2 S) + 1/2}, \quad (8)$$

and used instead of the normal value for ω in Eq. (3).

Shallow Water Simulation: Shallow water, or St. Venant, equations are usually used to simulate waves whose wavelength is similar to the overall water height. In this case the wave propagation speed is constant for all amplitudes. Deep water waves on the other hand are dispersive, which

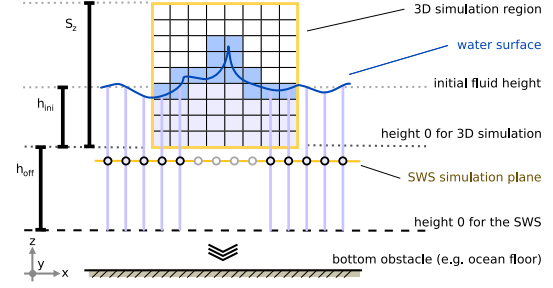


Figure 2: This picture gives an overview of our hybrid simulation method. The full three-dimensional fluid flow is solved in a given region of interest (illustrated by a 2D rectangle), and coupled to a two-dimensional shallow water simulation (shown as a 1D line in the picture).

means that the wave propagation speed depends on their amplitude. By using a shallow water model we make the assumption that for the limited range of amplitudes generated by the three dimensional simulation the wave propagation speed is the same. The advantage of a shallow water simulation is a full flow field for the water surface, that can produce vortices or handle e.g. flowing rivers.

Shallow water simulations (SWS) can likewise be performed using the LBM. In this case, instead of considering the fluid pressure, a height value is computed for each cell. Overall, the algorithm is very similar to the basic algorithm described above – both the streaming step and Eq.3 for relaxation towards the equilibrium are still valid. The equilibrium DFs to be used with Eq.3, however, are calculated differently. Furthermore, as the fluid surface is only two-dimensional, we use the *D2Q9* LBM model with nine velocities. To distinguish the DFs of the shallow water simulation from those of the three dimensional free surface simulation, we refer to them as g_l in the following. The fluid height h and the fluid velocity for the shallow water simulation are calculated as:

$$h = \sum_{l=1}^9 g_l \quad \mathbf{v} = \frac{1}{h} \sum_{l=1}^9 \mathbf{e}_l g_l. \quad (9)$$

In contrast to the 3D LBM model, the velocity computation of the SWS requires a division by the height, as shown in Eq.9. With height and velocity, the equilibrium DFs are computed as

$$g_0^{eq}(h, \mathbf{v}) = h \left[1 - \frac{5}{6} Gh - \frac{2}{3} \mathbf{v}^2 \right], \quad (10)$$

and

$$g_l^{eq}(h, \mathbf{v}) = w_l h \left[\frac{1}{6} gh + \frac{1}{3} \mathbf{e}_l \cdot \mathbf{v} + \frac{1}{2} (\mathbf{e}_l \cdot \mathbf{v})^2 - \frac{1}{6} \mathbf{v}^2 \right], \quad (11)$$

for $l = 1..9$. Here G is the gravity force, normal to the two-dimensional plane of the SWS, and the weights w_l have

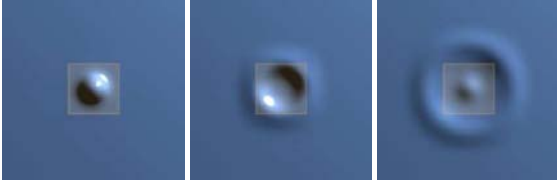


Figure 3: Wave propagation of a hemispherical drop on a flat surface using our algorithm. The 3D simulation region is highlighted in the middle.

the values $w_l = 1/18$ for $l = 2, \dots, 5$, and $w_l = 1/36$ for $l = 6, \dots, 9$. An in depth description of shallow water LBM can be found in e.g. [Zho04]. To establish a fixed height of the SWS boundary, we set the cells there to have the equilibrium DFs for the initial height and a zero velocity.

As this shallow water solver is similar to a basic LBM solver, the Smagorinsky turbulence model given by Eq.7 and Eq.8 can likewise be used to increase stability. The only difference is that $\Pi_{\alpha,\beta}$ is now computed as a sum over the nine DFs of a SWS cell with Eq. (6). To ensure stability for varying velocities, we furthermore apply the adaptive time stepping as described in [TPR*06] to both simulations.

3. Hybrid 2D/3D Simulation

An overview of our hybrid simulation approach is given in Figure (2). Both algorithms have been parametrized to solve the same fluid simulation problem, and are then coupled at a interface region. In the following we will assume that the SWS is performed in the xy -plane, and the gravitational force acts in the direction of the negative z -axis. There is an inherent difference between the two simulation approaches that has to be overcome: the derivation of the SWS assumes a depth averaged velocity and has a coupling between fluid height and velocity. The 3D simulation, on the other hand, can have a velocity varying along the z -axis, and has boundary conditions (see below) that makes it independent of the initial height of the fluid surface h_{ini} . In order to be able to couple both simulations, we have developed the following parametrization procedure for the SWS. It ensures that the SWS and the 3D simulation have the same wave propagation speed.

We offset the SWS by a constant height h_{off} , as shown in Figure (2). Here S_z is the height of the 3D domain in cells. In combination with the gravity the height offset h_{off} determines the wave propagation speed, and is set according to the average wave height generated by the 3D simulation h_{avg} . Assuming a common trochoid wave shape, we set h_{off} to be half of the expected wave length, thus $h_{off} = \pi h_{avg}$. For the examples shown in the following we have used a value of $h_{avg} = 1/2 h_{ini}$. Now the gravity force of the SWS has to be scaled according to the height offset. This is done by

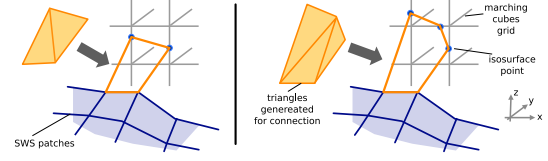


Figure 4: The two cases that need to be distinguished to generate a closed surface mesh for the 3D and shallow water simulations.

	G	N	Δt	\mathbf{v}
G	2	1	$\sqrt{2}$	$\sqrt{2}$
N	1	2	e	$\sqrt{2}$

Table 1: Behavior of the SWS upon parameter change. Here N is the number of SWS cells along the x - or y -axis.

examining the behaviour of the SWS properties. Given an arbitrary simulation setup, the properties of the fluid change by a factor given in Table (1), when the value of the parameter in the first column is multiplied by 2. Thus, given the initial SWS fluid height and $n = \log_2(h_{off} + h_{ini})$ we set the SWS gravity G to match the given offset for the simulation resolution. It is computed from the z -component of the 3D gravity \mathbf{g}_z as:

$$G = \mathbf{g}_z \cdot \left(\frac{e}{2}\right)^{-n}. \quad (12)$$

The 3D gravity is given by the physical value, usually $G' = 9.81 [m/s^2]$, as $\mathbf{g} = (0, 0, G' \cdot \Delta t^2 / \Delta x)$. The initial time step size Δt is set according to the maximum fluid or moving obstacle speed in the simulation setup. Now, when transferring velocities in the xy -plane between the simulations, the influence of the offset and gravity scale have to be removed. According to Table (1) this is accomplished by

$$\mathbf{v}_{x,y} = s_u \mathbf{u}_{x,y}, \text{ with } s_u = \frac{\sqrt{2}^n}{\sqrt{S_g} 1/n}. \quad (13)$$

3D to 2D Coupling: Here, we determine the height of the fluid at a position within the 3D simulation region by searching for the first interface cell. We start at the cell with grid position $(i, j, 0)$, and assume a planar fluid surface. Hence, the fluid height is computed for the first interface cell at (i, j, k_H) with

$$H(i, j) = k_H + \varepsilon(i, j, k_H) + h_{off}. \quad (14)$$

The velocity at the water surface \mathbf{u}_{surf} is given by the interface cell of the 3D simulation. To transfer the information from the 3D simulation to the SWS a cell at (i, j) , shown as a circle marked with X in Figure (5), is initialized with the

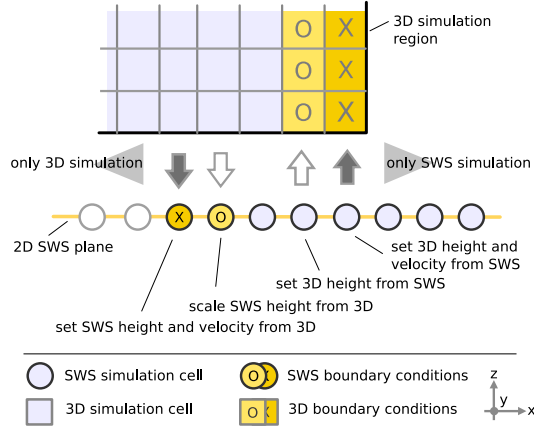


Figure 5: Detail of the double layer boundary conditions in the overlapping interface region.

equilibrium DFs:

$$g_l = g_l^{eq}(H(i, j), s_u \mathbf{u}_{surf}) . \quad (15)$$

The cells where height and velocity are set with Eq. (15) represent the inner boundary for the SWS. Further inwards we perform the full 3D simulation, thus the cells of the SWS do not have to be updated in this region, as their values are never used. To ensure a transfer with as few disturbances as possible, we use a double layered transfer. Thus, we use a second type of boundary condition for the region of SWS cells directly outwards of the boundary cells described above (circle marked with O in Figure (5)). For the cells that are updated according to Eq. (15) all DFs are reset each time step, while for the second boundary layer we only rescale the existing DFs to match the required fluid height:

$$g_l^* = g_l \cdot \frac{H(i, j)}{h(i, j)} . \quad (16)$$

The combination of these two boundary conditions ensures a correct transfer of both fluid surface height and velocity.

2D to 3D Coupling: The transfer from the SWS to the 3D simulation, is done by initializing the 3D cells to represent the SWS height and velocity. For a 3D cell at (i, j) marked with O in the upper part of Figure (5) we thus either remove cells from the simulation, if $H(i, j) > h(i, j)$, or otherwise add new ones. To correctly initialize the new cells we directly use the velocity of the SWS. The z coordinate of the velocity is calculated from the SWS fluid heights of the current and previous timestep.

For the outer layer of the 3D simulation (square cells marked with X in Figure (5)) we use velocity boundary conditions with a fixed pressure. For the LBM the pressure of a cell with height k in the domain is given by:

$$\rho_k = 1.0 + (h_{ini} - k) \cdot g_z \cdot -3\omega . \quad (17)$$

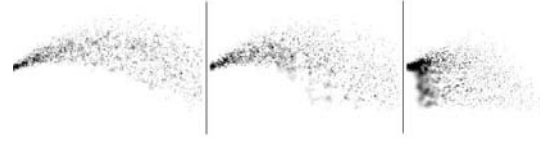


Figure 6: Effect of the drag model for different size scales. The simulations were parametrized to represent scales of 10cm, 1m and 10m, from left to right. The smaller particles, are slowed down, and cause mist below the outflow.

The pressure thus increases further down in the grid, with a gradient that depends on the relaxation time ω . The velocity can again be taken directly from the SWS, as described above. Note that it is in this case not necessary to scale the SWS velocities, as we set the whole height of the 3D simulation. These boundary conditions, however, do not ensure the full propagation of arbitrary waves generated in the SWS region, as this would require an additional wave profile initialization. Although these boundary conditions do not enforce mass conservation for the transfer, this is not problematic as the overall height of the fluid height is kept at the initial value by the SWS with its outer boundary conditions.

The depth of the overlapping region for the two simulations is variable, but we have found a distance of one eighth of the 3D domain size to yield good results. A validation run is shown in Figure (3). The circular wave retains its shape while it is transferred from the highlighted 3D region to the SWS region. The in- and outflow at the 3D domain boundary furthermore causes no disturbances of the flow field. A pure SWS simulation would not have been able to resolve the drop forming in the middle of the 3D region, visible in the right picture of Figure (3).

Surface Generation: To generate a mesh from the 3D simulation, we use the marching cubes algorithm. A triangulation of the SWS surface is easily computed by constructing patches between four adjacent SWS cells that have x and y coordinates according to their grid position, and a z coordinate given by $h(x, y)$. At the 3D domain boundary we leave out the first row of SWS cells, and construct triangles to connect the 3D mesh to the SWS patches. If both points of a marching cubes cell lie on its z-edges, this is sufficient to ensure a closed mesh. For all other cases, we also have to connect triangles to the points above or below the cell at the surface, as shown in Figure (4). In rare cases, e.g. when a drop directly hits the connection line, this technique will not result in a closed mesh. For our tests we have, however, not encountered any visible artifacts. In the interface region, where we have full information from both simulations, we linearly blend the fluid surface heights, to achieve a smooth transition from one type of simulation to the other. As the mesh generated from the fluid fractions already requires smoothing, we also perform a smoothing of the interface region to prevent any artifacts from misaligned normals.



Figure 7: A stream of water hits a rock and a water surface In the rightmost picture the interface between 2D and 3D region is highlighted.

4. Animation of Drops

For the animation of drops in our simulations we make use of methods developed for dispersed gas-liquid flows. In the following we will describe our model for the animation and generation of water drops. Each drop is described by its position \mathbf{x} , velocity \mathbf{w} and radius r . We assume that the drops are small enough to remain spherical due to surface tension. Thus using the density of water ρ_W the mass of a drop is given by its volume:

$$m_P = \rho_W \frac{4}{3} \pi r^3 . \quad (18)$$

Generation: To generate particles in our simulation we make use of the turbulence model explained in Section (2). As it already determines how many unresolved flow features a given cell has, we use it to compute a particle generation probability for cells at the fluid interface. We compute this probability from the absolute value of the Reynolds stress tensor given by Eq. (6) and the physical speed $\mathbf{u}' = \mathbf{u} \Delta x / \Delta t$. The stress tensor usually takes values of ca. $P_m = 10^{-2}$ for regions with significant unresolved detail independent of the actual grid resolution. We assume that the range of velocities, where the pressure of the surrounding air causes instabilities that lead to drops at the surface, is similar to that of the drop terminal velocities, which motivates the following probability function:

$$P_D = P_{ab} * (\mathbf{u}')^2 \quad \text{with} \quad P_{ab} = \sqrt{\Pi_{\alpha,\beta} \Pi_{\alpha,\beta}} . \quad (19)$$

Thus for a high physical speed of $10[m/s]$ and significant unresolved flow details this function will result in a drop generation probability close to one. Note that the calculation of the Reynolds stress tensor is especially easy for LBM, as it is computed locally from the derivative information contained in the non-equilibrium parts of the DFs (see Eq. (6)). For other types of solvers, this computation will require access to neighboring grid cells to compute the derivatives. Upon creation we initialize the drop velocity by the fluid velocity and a randomized normal offset to avoid immediate collision with the fluid surface.

Animation: For each LBM step, we simply update the particle position using their velocities and the LBM time step length:

$$\mathbf{x}(t + \Delta t) = \mathbf{x}(t) + \Delta t \mathbf{w} . \quad (20)$$

To update the particle velocities, we compute the balance of the forces acting upon it:

$$m_P \frac{d\mathbf{w}}{dt} = F_G + F_D . \quad (21)$$

where F_G is the force due to gravity and F_D is the drag force caused by the drop of water moving through the air. In contrast to the dispersed flow simulations mentioned above, we thus ignore any lift the drops might experience as well as other forces that would e.g. be caused by the density gradient in the air. The lift is proportional to the ratio between the involved fluids, which is close to zero for air and water. Likewise, we assume a density gradient in the air very close to zero. F_G is directly computed from gravity and particle mass as $F_G = m_P \mathbf{g}$, while the computation of the drag force requires more effort. The movement of water drops through the air has been studied in depth for meteorological purposes, see e.g. [PK97]. From these studies it is known, that rain drops usually have a size less than $4.5mm$. Above this size they will start to deform during their movement and eventually break apart due to the high forces from the air in comparison to the surface tension. It was furthermore measured, that these large drops have a terminal velocity of up to $w_{t1} = 9m/s$, while smaller drops of with e.g. $r = 0.5mm$ only accelerate to ca. $w_{t2} = 2m/s$. Given a coefficient of drag C_D , the drag force acting upon a particle is calculated as:

$$F_D = \frac{C_D}{2} \rho_L \pi r^2 \mathbf{w}_{rel} |\mathbf{w}_{rel}| , \quad (22)$$

where \mathbf{w}_{rel} is the relative velocity of the particle. It is computed from the velocity of the air \mathbf{w}_A by $\mathbf{w}_{rel} = \mathbf{w}_A - \mathbf{w}$. As we do not explicitly simulate the gas phase, we usually set $\mathbf{w}_A = 0$. Other values could be used to simulate the effect of wind. For the drag coefficient there are various approximations for different regimes of turbulence. As the case of a larger spherical rain drop at terminal velocity is already turbulent ($Re > 1000$), and the approximations are computationally very expensive, we fit the computation of the drag coefficient to yield values in the required range. We thus require that the drag force and gravity acceleration balance for the drops at their respective terminal velocities. Assuming a linear change for both parameters, we compute the drag force with a bilinear interpolation:

$$C_D = \frac{|\mathbf{w}_{rel}|}{w_{t2} + (w_{t1} - w_{t2})w_r} \left(\frac{1}{2} + \frac{1}{2}w_r \right) . \quad (23)$$

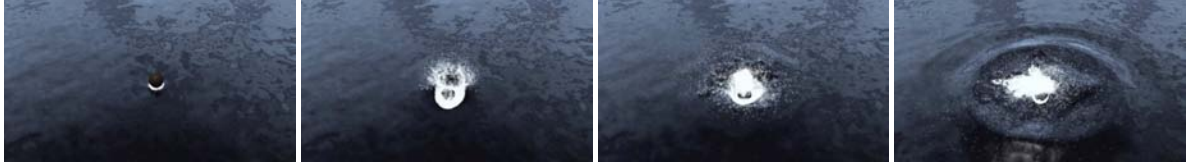


Figure 8: Animation of a spherical object hitting the water surface.

with $w_r = (r_1 - r)/(r_1 - r_2)$. For our simulations we have limited the size of the drops to the range of $r_1 = 0.005m$ to $r_2 = 0.0005m$. After the computation of F_G and F_D we update the velocity according to Eq. (21) with an Euler-step:

$$\mathbf{w}(t + \Delta t) = \mathbf{w}(t) + (F_G + F_D)m_p\Delta t. \quad (24)$$

Additional effects: To actually cause a disintegration of a thin fluid sheet into drops, we randomly choose a size r_D in the given range, and subtract the mass of the drop from the interface cell where it was generated. For simulations representing a large scale, this could result in huge numbers of particles – for these cases we subtract a multiple of the drop mass from the cell, and display the drop as a correspondingly larger transparent particle, thus representing multiple drops of similar size. Once the drop hits the fluid surface again, we add the mass that was subtracted before. Here, similar to [TFK*03], we trace the particle on the fluid surface to give the impression of foam. Figure (6) shows examples of the drag force influence for different scales. For the larger test cases the higher velocities result in higher drag forces, resulting in a noticeable slowdown of the smaller drops.

Another effect that cannot be directly simulated with the algorithm explained in Section (2), is that of instabilities caused purely by the relative physical velocity of the fluid \mathbf{u}'_{rel} , as the air is not simulated as a fluid itself. Thus, in order to cause these instabilities, we use a simple approximation, and manually add the following term

$$f_i = f_i + (P_m - P_{ab}) \cdot w_i \frac{\mathbf{u}'_{rel}}{50} \cdot \mathbf{e}_i, \quad (25)$$

for cells, with $P_{ab} < P_m$ that do not generate particles.

5. Results

All results shown in the following were created using a physical viscosity of water $\nu_W = 1 \cdot 10^{-6}$. To enhance the realism, we add a texture to the water surface, giving the impression of smaller chaotic waves. A test case of our simulation method is shown in Figure (7). A stream of water hits a rock and a water surface. The created waves spread outwards without a visible border between the SWS and the 3D simulation. Simulation resolutions and times can be found in Table (2). A test case that demonstrates the capabilities our drop model is shown in Figure (8). A spherical object is dropped into a fluid surface. The drop model, with up to

	3D	SWS	Simulation	Size
Figure (7)	120 ³	480 ²	14.6 s	0.2 m
Figure (8)	120 · 120 · 200	480 ²	34 s	10 m
Figure (9)	120 ³	960 ²	79.6 s	2 m

Table 2: This table shows grid resolutions together with average simulation times per frame (measured with an 2.2 GHz Opteron CPU). The size value is the physical length of a side of the 3D domain used for parametrization of the simulations.

50000 drops at a single time step, enhances the impression of a large simulation scale.

Given a working hybrid simulator, only small changes are necessary to achieve animations such as shown in Figure (9). Here we move the 3D domain according to the position of an object in the xy-plane. For each movement of the domain by Δx we copy the values stored in the grid by one in the desired direction. During the next step, the boundary regions of both simulations will again be correctly initialized for the boundary conditions. In this example the foam particles on the water surface clearly visualize the flow field in the shallow water region.

For a simulation run with a relatively large SWS domain, such as shown in Figure (9), we have measured the workload distribution between the different parts of the algorithm. In this case, ca. 68.8% are spent on the simulation of the 3D region. The 2D region, covering a 35 times larger area, requires 24.9% of the time, while ca. 2.6% are spent on the coupling of both simulations. The remaining 3.7% were spent on drop calculation, surface mesh generation and initialization. The update of an SWS cell is on average three times faster than the update of a 3D cell.

6. Conclusions and Outlook

We have presented an algorithm to efficiently simulate large scale open water fluids. This is achieved by coupling a shallow water simulator with a 3D free surface simulation. Our approach to model the formation and movement of drops allows us to add detail to relatively coarse fluid simulations. In

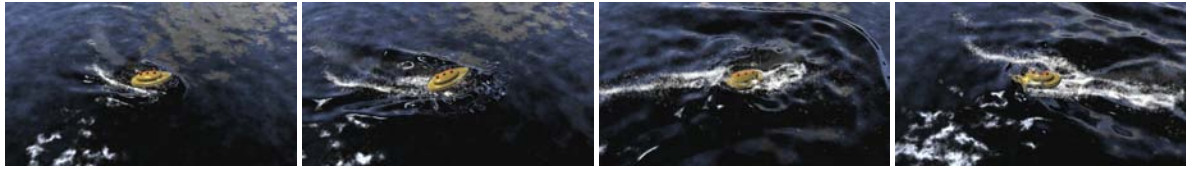


Figure 9: Pictures from an animation with a moving 3D domain.

conclusion, animations of large water surfaces, that would require huge amounts of memory and computational time with a conventional full 3D simulation, can be calculated within a matter of hours. The chance to arbitrarily move the 3D domain within the 2D region furthermore increases the flexibility of our approach.

As a future extension, we plan to increase the overall computational speed by making use of adaptive grids and parallelization for machines with multiple CPUs. For an intermediate scale, we are working on applying the methods from SPH to the generated drop particles, to accurately capture effects such as coalescence. It would furthermore be interesting to add a model for the generation of drops in the SWS region as well, or couple it with an FFT solver for ocean waves [Tes04]. An easy way to further speed up the computations would be to reduce the SWS resolution by an integer factor, and interpolate the values at its boundary. The problem of a fixed wave propagation speed, on the other hand, could be alleviated by overlaying multiple shallow water simulations with different parametrizations. Finally, the method could be used to couple multiple regions of three dimensional computation in one large water surface simulation. Given enough computational resources in combination with low grid resolutions, this could be used to simulate interactive environments with large water surfaces e.g. for virtual reality applications.

7. Acknowledgements

This research is funded by the DFG Graduate College GRK-244 *3-D Image Analysis and Synthesis*. We would also like to thank Vivek Buwa, Christian Feichtinger and Klaus Iglberger for helpful discussions.

References

- [BGD05] BUWA V., GERLACH D., DURST F.: Regimes of bubble formation on submerged orifices. *Phys. Rev. Letters* (April 2005). 2
- [Del01] DELLAR P. J.: Non-hydrodynamic modes and a priori construction of shallow water lattice Boltzmann equations. *Phys. Rev. E* 65 (2001). 2
- [DKvS99] DELNOI E., KUIPERS J. A. M., VAN SWAAIJ W. P. M.: A three-dimensional CFG model for gas-liquid bubble columns. *Chemical Engineering Science* 54 (1999). 2
- [FAM099] FEDKIW R. P., ASLAM T., MERRIMAN B., OSHER S.: A non-oscillatory Eulerian approach to interfaces in multicomponent flows. *J. of Comp. Phys.* 152 (1999), 457–492. 1
- [FdH*87] FRISCH U., D’HUMIÈRES D., HASSLACHER B., LALLEMAND P., POMEAU Y., RIVERT J.-P.: Lattice Gas Hydrodynamics in Two and Three Dimensions. *Complex Systems* 1 (1987), 649–707. 2
- [FM96] FOSTER N., METAXAS D.: Realistic Animation of Liquids. *Graphical Models and Image Processing* 58 (1996). 2
- [HHL*05] HAGEN T. R., HJELMERVIK J. M., LIE K.-A., NATVIG J. R., HENRIKSEN M. O.: Visual simulation of shallow-water waves. *Simulation Modelling Practice and Theory* 13 (2005). 2
- [HL97] HE X., LUO L.-S.: A Priori Derivation of Lattice Boltzmann Equation. *Phys. Rev. E* 55 (1997), R6333–R6336. 2
- [HNC02] HINSINGER D., NEYRET F., CANI M.-P.: Interactive Animation of Ocean Waves. 2
- [HSCD96] HOU S., STERLING J. D., CHEN S., DOOLEN G.: A Lattice Boltzmann Subgrid Model for High Reynolds Number Flow. *Fields Institute Communications* 6 (1996), 151–166. 3
- [IGLF06] IRVING G., GUENDELMAN E., LOSASSO F., FEDKIW R.: Efficient Simulation of Large Bodies of Water by Coupling Two and Three Dimensional Techniques. *ACM Transactions on Graphics* 25 (2006). 1, 2
- [KAG*05] KEISER R., ADAMS B., GASSER D., BAZZI P., DUTRE P., GROSS M.: A Unified Lagrangian Approach to Solid-Fluid Animation. *Proceedings of the 2005 Eurographics Symposium on Point-Based Graphics* (2005). 2
- [KM90] KASS M., MILLER G.: Rapid, Stable Fluid Dynamics for Computer Graphics. *ACM Trans. Graph.* 24, 4 (1990), 49–55. 2
- [KPR*05] KÖRNER C., POHL T., RÜDE U., THÜREY N., ZEISER T.: Parallel Lattice Boltzmann Methods for CFD Applications. In *Numerical Solution of Partial Differential Equations on Parallel Computers*, Bruaset A., Tveito A., (Eds.), vol. 51 of *LNCSE*. Springer, 2005, pp. 439–465. 3
- [LGF04] LOSASSO F., GIBOU F., FEDKIW R.: Simulating Water and Smoke With an Octree Data Structure. *ACM Trans. Graph.* 23, 3 (2004), 457–462. 1
- [MMS04] MIHALEF V., METAXAS D., SUSSMAN M.: Animation and Control of Breaking Waves. *Proceedings of the 2004 ACM SIGGRAPH/Eurographics Symposium on Computer Animation* (2004), 315–324. 2
- [PK97] PRUPPACHER H. R., KLETT J. D.: *Microphysics of Clouds and Precipitation*. Springer, 1997. 6
- [RNGF03] RASMUSSEN N., NGUYEN D. Q., GEIGER W., FEDKIW R.: Smoke simulation for large scale phenomena. *ACM Trans. Graph.* 22, 3 (2003), 703–707. 2
- [Sma63] SMAGORINSKY J.: General circulation experiments with the primitive equations. *Mon. Wea. Rev.* 91 (1963), 99–164. 3
- [Sta99] STAM J.: Stable Fluids. *Proceedings of the 26th conference on Computer graphics and interactive techniques* (1999), 121–128. 1
- [Tes04] TESSENDORF J.: Simulating Ocean Surfaces. *SIGGRAPH 2004 Course Notes* 31 (2004). 2, 8
- [TFK*03] TAKAHASHI T., FUJII H., KUNIMATSU A., HIWADA K., SAITO T., TANAKA K., UEKI H.: Realistic Animation of Fluid with Splash and Foam. *Computer Graphics Forum* 22 (3) (2003). 2, 7
- [TPR*06] THÜREY N., POHL T., RÜDE U., OECHSNER M., KÖRNER C.: Optimization and Stabilization of LBM Free Surface Flow Simulations using Adaptive Parameterization. *Computers and Fluids* 35 (8-9) (September–November 2006), 934–939. 4
- [TR04] THÜREY N., RÜDE U.: Free Surface Lattice-Boltzmann fluid simulations with and without level sets. 199–208. Workshop on Vision, Modelling, and Visualization VMV. 2
- [WZF*03] WEI X., ZHAO Y., FAN Z., LI W., YOAKUM-STOVER S., KAUFMAN A.: Natural phenomena: Blowing in the wind. *Proceedings of the 2003 ACM SIGGRAPH/Eurographics Symposium on Computer animation* (July 2003), 75–85. 2
- [Zho04] ZHOU J. G.: *Lattice Boltzmann Methods for Shallow Water Flows*. Springer, 2004. 4